En-Var in NWP (and ocean models)

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Basic problems of data assimilation

Filtering,

Interpolation

and

Balancing

Estimation of the characteristics of forecast errors (predictability) is important for all these three aspects of data assimilation !

HIRLAM first approach to use ensembles in 3D-Var and 4D-Var

- Use the ETKF algorithm for re-scaling of a 6h forecast ensemble to an analysis ensemble (estimation of the analysis error covariance).
- Use ensemble of 3h (4D-Var) or 6h (3D-Var) forecasts to estimate the background error covariance and blend it with the static background error covariance => Hybrid variational ensemble data assimilation, a step towards 4D-En-Var

Why 4D-En-Var?

- Avoid use of TL and AD models (low resolution TL and AD models have difficulties to scale on thousands of processors) => 4D-En-Var is cheaper than 4D-Var
- Utilizes 4D ensemble perturbations based on the non-linear model.
- Easy to implement with existing 4D-Var Hybrid
- 4D-En-Var in its simplest form is similar to 4D-En-KF. 4D-En-Var has possibilities to treat non-linearities better (outer loops); Easy to add 3D-Var FGAT background error constraint.



Incremental 4D-Var



Figure 3: Statistical, incremental, 4D-Var approximates entire PDF by a Gaussian. The 4D analysis increment is a trajectory of the PF model, optionally augmented by a model error correction term.

From Lorenc (2011)

4D-En-Var



Figure 6: A schematic diagram of 4D-En-Var, for comparison with figure 3. The 4D analysis is a localised linear combination of model trajectories – it is not itself a model trajectory.

From Lorenc (2011)

Incremental 4D-Var

Cost function minimized with respect to the increment $\delta \mathbf{X}$:

$$J = J_b + J_o = \frac{1}{2} (\delta \mathbf{X})^T \mathbf{B}^{-1} \delta \mathbf{X} + \frac{1}{2} \sum_{t_k=t_0}^{t_K} (\mathbf{H}_k \mathbf{M}_k \delta \mathbf{X} - \mathbf{d}_k)^T \mathbf{R}_k^{-1} (\mathbf{H}_k \mathbf{M}_k \delta \mathbf{X} - \mathbf{d}_k)$$

B the background error covariance $t_k = t_0, ..., t_K$ the data assimilation time window $\mathbf{d}_k = \mathbf{y}_k - H_k(M_k(\mathbf{x}_b))$ the innovations; \mathbf{y}_k the vector of observations at time t_k \mathbf{x}_b the model background state valid at time t_0 $M_k(.)$ the non-linear model; \mathbf{M}_k the corresponding tangent linear model $H_k(.)$ is the non-linear observation operator; \mathbf{H}_k the linearized observation operator \mathbf{R}_k is the observation error covariance

Introduce a pre-conditioning matrix U such that $\mathbf{B} = \mathbf{U}\mathbf{U}^T$, $\delta \mathbf{X} = \mathbf{U}\chi$

The cost function to be minimized and its gradient with respect to the assimilation control variable χ are given by:

$$J = J_b + J_o = \frac{1}{2}\chi^T \chi + \frac{1}{2} \sum_{t_k=t_0}^{t_K} (\mathbf{H}_k \mathbf{M}_k \mathbf{U}\chi - \mathbf{d}_k)^T \mathbf{R}_k^{-1} (\mathbf{H}_k \mathbf{M}_k \mathbf{U}\chi - \mathbf{d}_k)$$

and

$$\nabla_{\chi} J = \chi + \sum_{t_k=t_0}^{t_K} \mathbf{U}^T \mathbf{M}_{\mathbf{k}}^{\mathbf{T}} \mathbf{H}_{\mathbf{k}}^{\mathbf{T}} \mathbf{R}_k^{-1} (\mathbf{H}_{\mathbf{k}} \mathbf{M}_{\mathbf{k}} \mathbf{U} \chi - \mathbf{d}_k)$$

4D-Ens-Var (no localization)

Replace the static error covariance **B** with a flow-dependent error covariance $\mathbf{B} \approx \mathbf{X}'_{\mathbf{b}}(\mathbf{X}'_{\mathbf{b}})^T$ estimated from an ensemble of background model states

 $\mathbf{X}_{\mathbf{b}}'$ is a matrix whose columns are the normalized deviations of the ensemble background states from their mean:

$$\mathbf{X}_{\mathbf{b}}' = \frac{1}{\sqrt{N-1}} (\mathbf{X}_{\mathbf{b}\mathbf{1}} - \overline{\mathbf{X}_{\mathbf{b}}}, \dots, \mathbf{X}_{\mathbf{b}\mathbf{N}} - \overline{\mathbf{X}_{\mathbf{b}}})$$

N is the number of ensemble members.

Apply $\mathbf{X}_{\mathbf{b}}'$ for the pre-conditioning $\delta \mathbf{X} = \mathbf{X}_{\mathbf{b}}' \boldsymbol{\chi}$

Use an ensemble of non-linear model integrations over the data assimilation window and apply the following approximation in the observation constraint part of the cost function:

$$\mathbf{H}_{\mathbf{k}}\mathbf{M}_{\mathbf{k}}\mathbf{X}_{\mathbf{b}}' \approx \frac{1}{\sqrt{N-1}} (\mathbf{H}_{\mathbf{k}}(M_{k}(\mathbf{X}_{\mathbf{b}\mathbf{1}})) - M_{k}(\overline{\mathbf{X}_{\mathbf{b}}})), \dots, \mathbf{H}_{\mathbf{k}}(M_{k}(\mathbf{X}_{\mathbf{b}\mathbf{N}})) - M_{k}(\overline{\mathbf{X}_{\mathbf{b}}})))$$

The forward integration of the tangent linear model is replaced by the non-linear model trajectories.

The gradient of the cost function now becomes:

$$\nabla_{\chi} J = \chi + \sum_{t_k=t_0}^{t_K} (\mathbf{H}_k \mathbf{M}_k \mathbf{X}'_b)^T \mathbf{R}_k^{-1} (\mathbf{H}_k \mathbf{M}_k \mathbf{X}'_b \chi - \mathbf{d}_k)$$

Met Office

From Dale Barker (2011) Alpha Covariance Localization (Extreme example: 1 ob + 2 members!) • Single T observation (O-B, s_o=1K) at 50N, 150E, 500hPa.

To find the second seco

3D-Var



EnDA: No Localization

EXCRETE VIEW

EnDA: With localization



Crown Copyright 2011. Source: Met Office







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4D-En-Var (with localization)

En-KF: Covariance localization by an element-by-element multiplication (Schur product) with a full rank localization correlation matrix C:

 $\mathbf{B} \approx \mathbf{C} \circ \mathbf{B}_{ens} = \mathbf{C} \circ \mathbf{X}'_{\mathbf{b}} (\mathbf{X}'_{\mathbf{b}})^T$

C is constructed such that the final covariances $\mathbf{B} = \mathbf{C} \circ \mathbf{B}_{ens}$ will be zero over distances longer than a pre-defined localization length scale.

For 4D-Ens-Var the pre-conditioning can be done with the matrix $\mathbf{P}'_{\mathbf{b}}$ given by

$$\mathbf{P'_b} = (\mathbf{C'} \circ \mathbf{X'_{b1}}, \mathbf{C'} \circ \mathbf{X'_{b2}},, \mathbf{C'} \circ \mathbf{X'_{bN}})$$

 $\mathbf{C}'\mathbf{C}'^T = C$

 $\mathbf{X}'_{\mathbf{bk}}$ is an n-column matrix with every column being equal to the k^{th} column in $\mathbf{X}'_{\mathbf{b}}$ N is the number of ensemble members n is the model space dimension. \mathbf{C}' consist of eigenvectors of the correlation matrix \mathbf{C} .

- $J = \frac{1}{\beta_{3dvar}} J_{3dvar} + \frac{1}{\beta_{ens}} J_{ens} + J_o$ 4D-Var:
- $\delta x(t_k) = \mathbf{M}_k \delta x(t_0)$ 4D-Var Hybrid:
- $\delta x(t_k) = \mathbf{M}_{\mathbf{k}}(\delta x^{3dvar}(t_0) + \sum_{l=1}^{N} \alpha_l \delta x_l^{ens}(t_0))$
- 4D-En-Var:
- $\delta x(t_k) = \delta x^{3dvar}(t_0) + \sum_{l=1}^N \alpha_l \delta x_l^{ens}(t_k))$

Examples of ensemble spread (standard deviation) for temperature at model level 28 (~800 hPa)



Figure 12. Temperature level 28 spread (rms), 3dvar (top), 4dvar(bottom), before etkf re-scaling (left), after etkf re-scaling (right), 22

Before ETKF rescaling

After ETKF rescaling

Experiments over 17 January – 29 February 2008

4dvar_ref1: 4D-Var, 2 outer loops (6 h window, 20 iter. at 66 km and 40 iter. at 44 km incr. resol.), simple TL physics (vertical diffusion only), J_c DFI

4dvar_hybrid1: As 4dvar_ref1 with hybrid ensemble constraint, 20 members, ETKF perturb., 75% static and 25% ensemble variance, ens. perturbations inflated by a factor 4 in hybrid.

4DEnVar: 6 h window, 1 outer loop (60 iter. at 33 km incr. resol.). 50% static and 50% ens. variance, no ens. perturb. inflation, 3D-Var constraint in the middle of the window (<=> FGAT).



Model grid res. 11 km 40 levels 20 members

Example of single observation experiments with 4D-Var, 4D-Var Hybrid and 4D-En-Var

Background states



Position of simulated observation V500

Single observation assimilation increments

4D-Var



4D-En-Var



Hybrid impact on forecast verification scores – mean sea level pressure



---- 3D-Var; ---- 3D-Var hybrid ---- 4D-Var; ---- 4D-Var hybrid

48 stations Selection: EHGLAM Wind speed Period: 20080120-20080228 Statistics at 00 UTC At {00,12} + 12 24

Verification of wind speed and relative humidity from 6 weeks of parallel runs with HIRLAM 4D-Var, 4D-Var Hybrid and 4D-En-Var; 20 ensemble members.

Note that 4D-En-Var is much cheaper since TL and AD models are not needed (provided an ensemble exists!).



Sensitivity experiments

Is (further) inflation of ensemble perturbations needed? **4densvard (inflation 4.0) versus 4densvar (no inflation).** Answer: No!

Is static background error constraint needed ? 4densvar (50%static) versus 4densvarb (10% static) Answer: Yes (with 20 members only) !



41 stations Selection: EHGLAM Relative Humidity Period: 20080117-20080228 Statistics at 12 UTC At {00,12} + 12 24

"Correcting phase errors"

4dvar_hyb1: the 4DVAR hybrid (ETKF with 20 members) 4dvar_hyb2: the 4DVAR hybrid (ETKF with 40 members: 20 members: fc20080122_06+003 20 members: fc20080122_06+005)



Which ensemble generation technique is better?

ETKF or EDA



3DVAR-ETKF outperforms *3DVAR* and is slightly better than *3DVAR_EDA*

Dynamically consistent structures are important

EDA: analysis at 22 Jan 2008 12 UTC & mbr005



+000

Tue 22 Jan 2008 122 +00h - Tue 22 Jan 2008 122 valid Tue 22 Jan 2008 122



Tue 22 Jan 2008 122 +06h - Tue 22 Jan 2008 122 +06h valid Tue 22 Jan 2008 182





Tue 22 Jan 2008 122 +24h - Tue 22 Jan 2008 122 +24h valid Wed 23 Jan 2008 122

ETKF: analysis at 22 Jan 2008 12 UTC & mbr005



+000Tue 22 Jan 2008 122 + Tue 22 Jan 2008 122 -

Tue 22 Jan 2008 122 +008 valid Tue 22 Jan 2008 122









Is noise a potential problem for 4DEnVar (and ETKF re-scaling)?

• A weak digital filter constraint is applied in HIRLAM 4D-Var and HIRLAM 4D-Var Hybrid for the control forecast – no explicit initialization is applied.

• Do we need to apply initialization (incremental DFI) after ETKF re-scaling for ensemble members other than the control ?

• Do we need to apply initialization after 4D-En-Var, which is a hybrid of 3D-Var FGAT increment and localized ETKF non-linear model perturbations ?



- 4D-Var Hybrid Control is essentially noise-free

- 4dEnsVar control has a slightly incresed noise level

- Noise based on 4DEnsVar control increments and ETKF rescaling of ensemble perturbations adds up

Publications:

Bojarova, J., Gustafsson, N., Johansson, Å. and Vignes, O., 2010: The ETKF rescaling scheme in HIRLAM. *Tellus*, **63A**, 385-401.

Gustafsson, N., Bojarova, J. and Vignes, O., 2014: A hybrid variational ensemble data assimilation for the HIgh Resolution Limited Area Model (HIRLAM). *J. of Nonlin. Processes in Geophys.*.

Gustafsson, N. and Bojarova, J., 2014: 4-Dimensional Ensemble Variational (4D-En-Var) data assimilation for the HIgh Resolution Limited Area Model (HIRLAM). *J. of Nonlin. Processes in Geophys.*.

Verification of the HARMONIE AROME 2.5km forecasts for extreme weather event (from Xiaohua Yang (DMI) & Lisa Bengtsson et al (SMHI))

Radar data 31.08 00UTC - 12UTC



HARMONIE AROME + 30h (MetCoOp)





The HARMONIE AROME **is capable** in many cases to predict convective precipitation events (severe high impact weather events);

Stochastic nature of the convective phenomena should be taken into account both for verification and in post-processing (timing and location uncertainty);

The quality of the short-term forecasts in the operational runs is not satisfactory : **coupling strategy and data assimilation to be blamed**

Evolution of two random perturbations with structure of B-matrix covariance

Forecast length: +00h



Evolution of two random perturbations with structure of B-matrix covariance

Forecast length: +01h



Surface pressure increment **13 08 2012 04UTC**



Evolution of two random perturbations with structure of B-matrix **covariance**

Forecast length: +05h

Surface pressure increment 13 08 2012 08UTC





What structure functions say



Aliasing of high-order terms on $2\Delta x$, $3\Delta x$, $4\Delta x$, $5\Delta x$ waves Obvious 2Δx problem

The preliminary results using **cubic grid truncation** (Mariano Hortal implementation) show encouraging results : **increased numerical stability of the scheme and longer time stepping in the semi-lagrangian forward propagation**. Processes are resolved in the grid-point space and smoothed in the spectral space.





Average in time (25 cases)

Climatological structure functions (6 EDA based HarmonEPS perturbations; 06UTC + 12h)





Climatological structure functions (6 EDA based HarmonEPS perturbations; 06UTC + 12h)



What can we learn from this experiment:

1) We cannot come much further forward without flowdependent structure functions!=> homogeneity and isotropy assumption about the forecast error statistics do not hold for the convective scale phenomena;

2) **Small scales structures and noise is a dangerous combination** => Go for "cubic grid" truncation, possibly low-resolution orography; We need to rethink about initialisation on convective scales

3) **Ensembles have big potential for data assimilation on convective scales (processes driven by surface and PBL conditions)** => Go for Ensemble Variational techniques using convection permitting ensembles

Status of NWP data assimilation developments

- ECMWF: Global 4D-Var. Add model error and ensemble components. EDA for EPS – variances and now also correlations. EnKF option.
- NCEP: Global and LAM 3D-Var. Flow-dependent B=>Hybrids. 4D-En-Var.
- UKMO: Global and LAM 4D-Var. Hybrid En-Var => 4D-En-Var.
- M-France: Global 4D-Var. LAM 3D-Var. Ensemble B. 4D-En-Var
- DWD: Global 3D-Var. LAM nudging. Considers flow-dependent B globally and LAM EnKF.

- Japan: Global and LAM 4D-Var. Considers EnKF for the global model.
- Canada: 4D-Var for deterministic models, EnKF for EPS. Hybrids.4D-En-Var
- HIRLAM: 4D-Var. Hybrid En-Var. 4D-En-Var.
- ALADIN/HARMONIE: 3D-Var and development of 4D-Var. Ensemble components.
- University world: strong dominance for EnKF

3D-En-Var in an operational ocean model

- (Lars Axell, SMHI, personal communication) **Some basic features**
 - HIROMB or NEMO models
 - Minimization of a cost function
 - Quasi-static ensemble (multi-year model integration; use of ensemble members from the same season)
 - Localization function through EOFs and low resolution

Why 3D-En-Var when an Ensemble OI is available?

- Non-linear observation operators
- Step toward (3)4D-En-Var with forecast ensemble

Operational HIROMB domain:



EOF truncation of localization function at low grid resolution



Single observation experiments

Surface salinity:

Surface temperature:



FIG. 4. Result from assimilating six synthetic observations of (a) SSS and (b) SST. The observation locations are indicated with crosses (+).

Assimilation of real observations, Sea Ice Concentration

Innovations:

Analysis increments:



25 years reanalysis;

Salinity profiles;

Dependent observations



FIG. 12. Salinity profiles at station BY15, according to (a) observations, (b) a free run (without data assimilation), and (c) the reanalysis (with data assimilation). The data are sampled once a month.

25 years reanalysis;

verification against independent observations





FIG. 14. Bias and rms errors from two 25-year simulations, one free simulation ("No DA") and a reanalysis run ("With DA"). The panels show (a) near-surface salinity (0-50 m), (b) sub-surface salinity (50-500 m), (c) near-surface temperature (0-50 m), and (d) sub-surface temperature (50-500 m).

Concluding remarks

Adding ensemble information provides improvements to 3D-Var and 4D-Var (Hybrid Var Ens) (ds=10km)

Replacing the TL /AD models in 4D-Var with a 4D ensemble (4D-En-Var) provides results comparable to 4D-Var Hybrid

Several factors contribute to difficulties for 3D-Var assimilation at meso scale (ds=2.5km)

- assumptions on stationarity, homogeneity and isotropy are not valid

- no initialization

- ensemble information has potential to help, localization becomes an important issue